

Diagnosing Sources of U.S. Seasonal Forecast Skill

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Abstract

In this study we diagnose the sources for the contiguous U.S. seasonal forecast skill that are related to sea-surface temperature (SST) variations using a combination of dynamical and empirical methods. The dynamical methods include ensemble simulations with 4 atmospheric general circulation models (AGCM) forced by observed monthly global SSTs from 1950 to 1999, and ensemble AGCM experiments forced by idealized SST anomalies. The empirical methods involve a suite of reductions of the AGCM simulations. These include uni- and multi-variate regression models that encapsulate the simultaneous and 1-season lag linear connections between seasonal mean tropical SST anomalies and U.S. precipitation and surface air temperature.

Nearly all of the AGCM skill in U.S. precipitation and surface air temperature, arising from global SST influences, can be explained by a single degree of freedom in the tropical SST field --- that associated with the linear atmospheric signal of El Niño/Southern Oscillation (ENSO). The results support previous findings regarding the preeminence of ENSO as a U.S. skill source. Our diagnostic methods exposed another skill source that appeared to be of non-ENSO origins. In late fall, when the AGCM simulation skill of U.S. temperatures peaked in absolute value and in spatial coverage, the majority of that originated from SST variability in the sub-tropical west Pacific Ocean and the South China Sea. Hindcast experiments were performed for 1950-1999 that revealed most of the simulation skill of U.S. seasonal climate to be recoverable at 1-season lag.

The skill attributable to the AGCMs was shown to achieve parity with that attributable to empirical models derived purely from observational data. The diagnostics promote the interpretation that only limited advances in U.S. seasonal prediction skill should be expected from methods seeking to capitalize on sea surface predictors alone, and that advances which may occur in future decades could be readily masked by inherent multi-decadal fluctuations in skill of coupled ocean-atmosphere systems.

1. Introduction

Foremost among U.S. seasonal forecast skill sources are the state of tropical sea surface temperatures (SST), a relation stemming from the sensitivity of atmospheric stationary waves and storm tracks to tropical forcing (e.g., Hoskins and Karoly, 1981; Held et al. 1989). To date, the El Niño-Southern Oscillation (ENSO), that is typified by east tropical Pacific SST variations, is the single phenomena of the ocean-atmosphere system demonstrated to render U.S. seasonal forecasts skillful (e.g., Barnett and Prisendorfer 1987; Barnston 1994; Higgins et al. 2004). That skill is elevated in the winter and spring seasons, at which time it also exhibits a large spatial footprint over the U.S. due to the planetary-scale of ENSO-forced circulation variations.

The question we pose is whether ENSO constitutes the sole tropical SST source of U.S. seasonal forecast skill. To be sure, empirical and simulation studies have identified *sensitivities* to SST forcings outside the east tropical Pacific. For example, the influence of SST anomalies over a region north of the equator in the tropical west Pacific was highlighted by Palmer and Owen

(1986) as contributing to the severe U.S. winter of 1976-77. Empirical support for such a connection was provided by studies of the relation between atmospheric circulation patterns and proxies for tropical convection (Livezey and Mo 1987). The influence of SST anomalies over the warm pool regions of the Indian and western Pacific Oceans have also been shown to play roles in the U.S. climate, such as during 1997 and 1998 (Pegion et al. 2000), and during 1998-2002 when the U.S. experienced protracted warm and dry conditions (Kumar et al. 2001, Hoerling and Kumar 2003). Idealized modeling experiments, both linear dynamical models forced by tropical heating (e.g., Ting and Yu 1998; Chen and Newman 1998), and perpetual-mode general circulation models forced with idealized SSTs (Geisler et al. 1985; Barsugli and Sardeshmukh 2002), have confirmed robust Pacific-North American responses to forcings located outside the classic ENSO region of the east tropical Pacific.

Historical simulations of 20th Century climate using atmospheric general circulation models (AGCMs) subjected to observed monthly evolving global SSTs are now permitting diagnosis of sensitivity patterns under realistic conditions. The availability of multi-AGCMs, each run numerous times over an historical period, permits a more robust and reliable identification of atmospheric response patterns associated with tropical forcing than hitherto possible. Using a 46-member ensemble of 50-year experiments spanning 1950-1999 from 4 different AGCMs, Hoerling and Kumar (2002) subjected the monthly 500 hPa height fields of the multi-model ensemble to an empirical orthogonal function (EOF) analysis. These yielded evidence for at least 4 different characteristic atmospheric patterns associated with tropical SST variability. The leading one was the well-known ENSO teleconnection. However, the additional patterns accounted for up to 50% of the local boundary forced variance over parts of North America, and

appeared to be of non-ENSO tropical origin. It is unknown whether these additional sensitivities contribute to *skill* of U.S. surface temperature and precipitation predictions.

An ensemble of AGCM runs yields a wealth of information about the sensitivity of the atmosphere to observed SST anomalies. However, it is often not clear what the relationship between the prescribed SST and the GCM ensemble mean atmospheric response is. For example, further diagnosis is necessary to determine how many degrees of freedom in the SST field are actually involved in producing the simulated atmospheric signals. This is useful information, since it defines the relevant phase space of SST patterns that need to be predicted accurately for a seasonal forecast. Even if information can be obtained about the SST patterns from the AMIP runs (see e.g. Hoerling and Kumar, 2002), one is still left wondering how predictable those patterns are. Goddard and Mason (2002) have examined the predictability issue by running a complementary ensemble of GCM runs, but instead of the traditional AMIP approach, they use SSTs that are persisted (as a simple prediction) for one season. In this way they were able to investigate the predictability of the seasonal mean atmospheric state based on forecast SSTs rather than simultaneous SSTs. Here we attempt to address both the question of how many degrees of freedom are needed to reproduce the ensemble mean GCM response over the U.S., and the question of how predictable that response is, using an empirical statistical model trained on the 46 member multi-model GCM ensemble.

Our study combines AGCM and empirical methods to diagnose the simulation and hindcast skill of seasonal surface climate over the U.S. during 1950-1999. The data sets and methods are

presented in section 2. Section 3 presents results in which the seasonally varying skill of surface temperature and precipitation are described and diagnosed. In addition to the expected ENSO contribution, a non-ENSO source of skill is discovered. Section 4 focuses on a diagnosis of this additional source of U.S. skill. A summary and concluding comments are given in section 5.

2. Data and Method

a. Observations

Monthly observed U.S. surface temperature and rainfall data for 1950-1999 are available on a 2.5 ° grid based on the Global Historical Climate Network. Monthly sea surface temperature analyses for 1950-1999 come from the United Kingdom Meteorological Office (UKMO) Hadley Centre's Global Sea-Ice and SST (HadiSST) dataset (Rayner et al. 2003). The global SSTs are created using various techniques including reduced space optimal interpolation, and are available on a 1° grid.

b. Atmospheric climate simulations forced by global SSTs

The sea surface temperature role in climate variability is assessed using atmospheric GCMs forced with the specified, observed monthly variations in SSTs for 1950-1999. Multiple integrations are begun from different atmospheric initial conditions, but in which each ensemble member is subjected to identically specified sea surface conditions.

Four different AGCMs are used, consisting of a total of 46 simulations spanning the last half of the 20th Century. The models used are identical to those in Hoerling and Kumar (2002). Table 1 summarizes their spatial resolutions, and the reader is referred to the indicated references for

details on different dynamical cores and physical parameterizations used in each model. The atmospheric GCMs include 12 simulations using NCAR's Community Climate Model (CCM3; Kiehl et al. 1998), 12 using NCEP's Medium Range Forecast model (MRF9; Kumar et al. 1996), 10 with the European Centre-Hamburg Model (ECHAM3; Roeckner et al. 1992) and 12 simulations using the climate model of the Geophysical Fluids Dynamics Laboratory (GFDL; Broccoli and Manabe 1992).

Our analysis of the AGCMs is based on the combined 46-member ensemble. The ensemble mean seasonal anomaly for each of the four models is calculated relative to the respective models' climatology for each of the 600 overlapping seasons during 1950-1999. The ensemble mean anomaly of each model is standardized by its SST-boundary forced (external) variance, as described in more detail in Hoerling and Kumar (2002). The ensemble mean seasonal anomaly for each of the 4 GCMs is calculated, with anomalies computed relative to the respective model's 1950-99 climatology. Because these models possess different sensitivities to identical boundary forcing, we standardized the ensemble anomaly of each GCM by its own external variance, also calculated for the 1950-99 period. These are then averaged across the 4 models to yield a grand, standardized anomaly (consisting implicitly of the average anomaly for 46 discrete realizations). Actual anomalies are recovered by multiplying the standardized anomaly of each season by the 4-model averaged external variance. These standardized anomalies are then averaged across the four models in order to produce the multi-model, standardized anomaly. Actual anomalies are recovered by multiplying the standardized anomaly of each season by the four model averaged external variance. Horizontal and vertical resolutions of these models are listed in Table 1. In an analysis examining the impact of different ensemble mean methods, Byun (2002, personal

communication) compared the skill of this 46-member ensemble mean and the skill of another ensemble mean of the same 46-member but using a weighted average in which the weights were different for each member and determined by minimizing the difference between the model ensemble mean and the verification field. Byun's results indicated that for the 1950-99 50-year period the simulation skill of the weighted ensemble mean is virtually the same as the skill of the ensemble mean used in this study.

c. Atmospheric climate simulations forced by idealized SSTs

Simulations using a fixed idealized SST anomaly over the subtropical western Pacific are conducted, the pattern of which is discussed in section 4. This anomaly is specified and fixed throughout the seasonal cycle, and is added to the seasonally varying climatological SSTs. A total of 20 14-month simulations, beginning from randomly selected 1 November atmospheric initial conditions, were performed with NCAR's CCM3, one of the models included in our multi-model suite forced by realistic SSTs.

d. Empirical climate prediction model

The empirical model is based on uni- and multi-variate linear regression that relates a set of seasonally averaged fields of tropical (20°S-20°N) SST (the predictors) to a set of seasonally averaged model simulated fields of U.S. surface temperature and precipitation (the predictands) (see Fig. 1). Regression models are developed both for the simultaneous and 1-season lag relationships between predictor-predictand pairs. This allows us to estimate the relevant SST subspace both for the zero-lag (contemporaneous) GCM *simulation*, and the 1-season lag GCM *prediction*. Unlike the Goddard and Mason study, we do not rerun the GCM with predicted

SSTs, but instead use the lag-covariance relationships between the SST and the simulated atmospheric response from the AMIP integrations to estimate what the GCM prediction would be, given actual predicted SSTs. This approach assumes that the relationship between SSTs and the atmospheric response over the U.S. is linear. This hypothesis is tested in the next section by verifying that the empirical model can reproduce the GCM simulation skill over the U.S.. The empirical statistical model we use is very similar to others used for seasonal prediction (e.g., Barnston 1994), except for the fact that in our case the model is trained on the 46-member ensemble average of multi-model simulations described in section 2b, instead of observed atmospheric states. By training on the model ensemble mean, our tool describes the relation between atmospheric *signal* and tropical SST forcing, whereas similar tools trained on the brief, “single realization” of observed states necessarily predict a blend of signal and noise given a tropical SST state.

In the construction of the statistical model, the raw, seasonal predictor and predictand data are first pre-filtered using EOFs in a manner outlined by Barnett and Preisendorfer (1987) and as implemented by Barnston (1994). We retain N-EOFs in predictor and M-EOFs in predictand fields, where the truncation is determined by applying the formalism of North et al. (1982). For the tropical SSTs, N is chosen to be 5 for all 12 overlapping seasons. For U.S. precipitation, M is chosen to be 6 for all seasons, while M is chosen to be 9 for U.S. surface temperature. The North et al. (1982) analysis suggested that the EOF modes higher than a critical order are statistically indistinguishable. Therefore, we do not include higher order SST modes in our analysis and treat the higher order SST modes as noise. The EOF expansion coefficients (principal components) are then normalized by the square root of their eigenvalues, and the EOF

basis is calculated independently for each of 12-overlapping 3-month seasons. The cross-covariance matrix is calculated in EOF space, and the predicted EOF expansion coefficients are transformed back to a 2.5 degree latitude/longitude grid. A parallel set of uni-variate regression models is constructed to isolate the contribution of ENSO to the predictands of interest. The uni-variate regression uses only the 1st EOF (i.e. ENSO) mode of the tropical SST as a predictor, and 6 (9) EOF modes of the U.S. precipitation (surface temperature) as predictands. The uni-variate regression is also trained on the multi-model ensemble mean.

d. Verification methodology

We employ a deterministic measure of skill based upon the temporal correlation. Monthly mean surface air temperature and precipitation anomalies are obtained by subtracting the 30-year climatology of 1970-99 for each calendar month from the monthly mean values at each grid point over our domain (see Fig. 2). Three-month mean anomalies are then calculated from the monthly mean anomalies. The same method is applied to the output from the 46-member AGCM simulations. Verification is performed using the three-month mean anomalies.

All the uni- and multi-variate empirical model simulations and hindcasts for each 3-month season employ a cross-validation procedure in which the model is constructed from the covariance matrix of 49-years of seasonal data, excluding the target season for the simulation or hindcast. In this manner, a “new” empirical model is constructed for each year, and also for each season. The 1951-1999 skill of the empirical model simulations and hindcasts was then calculated in the way identical to that applied to the AGCM simulations. When calculating spatial averages of correlation skill for U.S. surface temperature, coastal gridpoints (indicated by “o” in Figure 2) are excluded because local temperature skill at those areas can be largely

determined by the prescribed SST in AGCM simulations.

The field significance of the correlation skill maps for each variable and for each of the twelve 3-month seasons is tested in a manner similar to Livezey and Chen (1983). For a given variable and season, Monte Carlo experiments of 3000 samples are first performed by randomly ordering the 50-year time sequence of observed fields and then calculating temporal correlations with the true 1950-99 observed time sequence. From these, we compute the probability distribution of the number of U.S. points in Fig. 2 having temporal correlations exceeding 90% local significance ($r > 0.25$). Then, the number of U.S. grid points (N_0) at 95% of the accumulated obtained distribution constitutes the minimum number necessary for passing field significance. In order that a correlation skill pattern be field significant, the actual number of locally significant grid-points (N) must exceed N_0 , i.e. the ratio N/N_0 must exceed 1.

3. Diagnosing SST skill sources

a. U. S. Precipitation

The annual cycle of simulation skill, and its sources, averaged over the U.S. is summarized in Fig. 3. Correlation skill of the AGCM simulations is low in all seasons (top), having a minor maximum during winter (thick solid line). Its primary source is tropical SSTs, as demonstrated by the ability of the multivariate regression model simulation skill (thin solid line) to fully recover that of the global SST forced AGCMs. Further, a single pattern of the tropical SST variations is able to explain this skill (dotted curve), namely the ENSO signature of the tropical east Pacific (not shown). The simulation skills are field significant in winter and spring, but not so in summer and early fall (bottom).

Inspection of the spatial patterns of skill reveals the univariate regression model to be superior to the AGCM simulation in select regions. Figure 4 shows the temporal anomaly correlation for the U.S. domain, and the plotted values are the truncated correlation skill scores ($\times 10$) at all points exceeding 90% significance. The lower right corner of each panel plots the ratio N/N_0 , which denotes field significance for values >1 . The six columns correspond to partially overlapping 3-month seasons, and the top three panels of each column display the AGCM simulation skill, the multivariate regression model skill, and the univariate regression model skill, respectively.

During January-February-March (JFM), significant skill covers nearly half the domain in the univariate model, and mimics the spatial footprint of ENSO's influence on U.S. precipitation (e.g. Ropelewski and Halpert 1989, Kiladis and Diaz 1989). This pattern is only weakly discernable in the AGCM simulation skill. Throughout all remaining seasons, it is found that the AGCM simulation skill is generally surpassed by a reduced-space representation of the relation between SST and model U.S. precipitation.

The lower panels of each column display the zero-lead hindcast skill using the multivariate model. These recover virtually all of the 50-yr averaged AGCM simulation skill. The uni-variate model often performs better than the multi-variate model. This is so because ENSO SSTs are the primary skill source, and the fact that such an ENSO pattern can itself be skillfully predicted at short lead times.

b. *U. S. Surface Temperature*

The SST source for U.S. surface temperature skill is found to be determined not by ENSO alone, in contrast to the pure ENSO-source that was diagnosed for U.S. precipitation skill. Figure 5

summarizes the area-averaged skill of the three simulation data sets, from which it is evident that the three simulation skills are comparable during winter and spring, but that the AGCM skill is greater than either multi- or univariate regressions in summer and fall. Note that the AGCM skill is field significant in all seasons (bottom), contrary to the empirical model skill which are not field significant from September-November (SON) through November-January (NDJ).

There is a distinct NDJ seasonal peak in simulation skill, at which time the U.S. averaged 50-yr mean skill exceeds 0.36. The field significance exceeding 99.9% is all the more remarkable when compared to the lack of field significance by either empirical model simulation. By contrast, the coincidence in skill of the three simulations during winter and spring indicates that ENSO's linear signal is the primary source. An ENSO SST source appears to impart little skill during other seasons as indicated by the large divergence between the AGCM and regression model skills at those times. These other SST sources could be of tropical origin, but are unaccounted for by our truncated multivariate model, or they could be of extratropical sea surface origin. It is also possible that non-linear processes associated with the atmospheric response to SST forcing are important contributors to the AGCM skill, relations also not represented in the regression models.

The spatial distribution of surface temperature skill, shown in Fig. 6, offers further insights on SST sources. There is little appreciable difference in skill patterns occurring among the three simulations during winter and spring --- the similarity with the univariate model skill confirms once again that the linear component of the ENSO-forced atmospheric signal is of primary consequence. On the other hand, the AGCM simulation skill is greater in both its coverage and

its absolute value than either regression model during summer and fall. The footprint differs from the linear ENSO contribution which, as illustrated by the univariate skill map, is confined to the southern U.S..

Shown in the lower panels of Fig. 6 are the hindcast skills for the multivariate model. These recover much of the simulation skill of U.S. surface temperature during winter/spring, consistent with the fact that the linear ENSO influence is of greatest relevance and that such SSTs are themselves skillfully predicted at zero-lead. In contrast, they fail to reproduce the extensive spatial coverage of the AGCM's simulation skill during the NDJ season.

4. A Subtropical West Pacific Skill Source

We first explore the SST structure associated with the AGCM's high skill over the northern U.S. during November-December-January. The green box in Fig. 7 keys on the region of high AGCM skill, and is the domain over which surface air temperatures are averaged in order to generate a 1950-1999 time series. The correlation of that time series with SSTs across the global oceans is shown by the shaded contours, and the analysis is performed using both the AGCM (top) and observed (bottom) surface air temperatures. Maximum correlations exceeding +0.5 occur with SSTs over the northwest Pacific Ocean and the South China Sea, and there is also some suggestion of an ENSO relation in so far as +0.3 correlations occur with over the tropical east Pacific. However, the fact that an ENSO-region correlation is very weak in the observations suggests it is unlikely an appreciable skill source. By contrast, the consistency between the simulated and observed high correlations over the Western Pacific suggests that SST variations there could be an important skill source for northern U.S. surface air temperature.

The coherency of their interannual fluctuations lends some support for a causal relation between subtropical west Pacific SSTs and the northern U.S. index region's temperatures. Figure 8 overlays the times series of the surface air temperature index and an index of SST averaged over the west Pacific box of Fig. 7. The former is calculated from both AGCM (Fig. 8, top) and observed (Fig. 8, bottom) surface air temperatures. For both, the correlation with the west Pacific SST index exceeds +0.5, with warm sea surface conditions accompanying warm northern U.S. surface temperatures. In addition to the tendency for interannual swings of each index to be aligned, a post-1970 warming trend in both ocean and land time series also contributes to the positive temporal correlations.

Shown in Fig. 9 is the temporal variation of our west Pacific SST index with SST elsewhere. The coherency of an SST signature that encompasses the South China Sea, Philippine Sea, and the subtropical west Pacific is evident. There is only a weak simultaneous relation with SSTs over the ENSO region, again suggesting that this west Pacific source of U.S. simulation skill in the AGCM is not an obvious proxy for ENSO's effect.

Independent evidence for a causal link between variations in west Pacific SSTs and U.S. surface temperatures is provided by the additional AGCM experiments forced exclusively with SSTs over that region. The idealized anomalous SST forcing is obtained from a composite of the SST pattern contrasting warm and cold events (when anomalous values exceeded one standard

deviation) in the northern U.S. temperature (Fig.8 top), and has similar structure as outlined by the irregular contour in Fig. 9, and encompasses the spatial domain of temporally coherent west Pacific SST variations. A 20-member ensemble of CCM3 simulations was conducted using both positive and negative polarities of the SST anomalies, with maximum amplitudes of 1° C. The NDJ surface temperature anomalies of the (warm-cold) experiments is shown in the lower panel of Fig. 10. Widespread warming covers North America, with a maximum signal over the northern U.S. and southern Canada. This response pattern is remarkably similar with the 1950-1999 correlation between an index of the west Pacific SSTs and both multi-model AGCM (Fig. 10, top) and observed (Fig. 10, middle) surface temperatures. Indicated hereby is that the statistical correlation does indeed describe a cause-effect relationship.

The implications of the above analyses is that a bivariate regression model, having west Pacific and tropical east Pacific SST predictors, should be able to recover the AGCM simulation skill during the NDJ season. The simulation skill of such a bivariate model is shown in Fig. 11 (middle panel). For comparison, the simulation skill of only a univariate model using the west Pacific SST as predictor is shown in the top panel of Fig. 11. Nearly all of the AGCM's simulation skill over the northern U.S. during NDJ (see Fig. 6) is accounted for by a west Pacific influence. The results of the bivariate model, which also includes the ENSO influence, almost completely recovers the AGCM simulation skill. These calculations support the argument that the high AGCM simulation skill for NDJ surface temperature over the northern US is attributed to the ocean-atmosphere interactions in the Western Pacific Ocean.

As a final analysis, we re-examine the hindcast skill of NDJ surface temperatures, but now using a 1-season lag bivariate regression model. The results are shown in the lower panel of Fig. 11, and it is apparent that the vast majority of simulation skill is retained in the prediction mode. This suggests that the west Pacific SST predictor is itself predictable and slowly evolving on the seasonal time scale. The new subtropical SST source bears some resemblance to the 3rd CCA SST mode in Barnston and Smith (1996). However, our analysis revealed that the SST variation in the subtropical Northwest Pacific Ocean is not a direct response to ENSO. In fact, the SST variation in the subtropical NW Pacific Ocean leads an out-of-phase ENSO episode by about one year (not shown).

5. Summary and discussion

a. Summary

The sea surface temperature origins for seasonal forecast skill of U.S. precipitation and surface air temperature have been diagnosed. The principal tool consisted of ensemble atmospheric GCM simulations that were subjected to the known monthly varying global SSTs from 1950-1999. The simulation skill of the resulting multi-model 46-member ensemble was evaluated, and then diagnosed using two empirical reductions of the full AGCM data. The first consisted of a multivariate linear regression model relating the simultaneous states of five leading tropical SST patterns to the AGCM's U.S. seasonal climate. The second was a univariate linear regression model whose predictor consisted of the single leading tropical SST pattern only --- ENSO.

Our diagnosis revealed that the AGCM simulation skill could be explained by-and-large from the

linear influence of tropical sea surface temperature variations. Furthermore, one degree of freedom in those tropical SSTs, namely ENSO, was the virtually exclusive interannual skill source originating from tropical oceans. In the case of U.S. seasonal precipitation, the univariate regression model that encapsulated the linear ENSO signal produced simulation skill that exceeded the AGCM simulation skill, though field significance was confined to the period late Fall–late Spring. One interpretation for the superior precipitation skill of the univariate model is that whereas sensitivity patterns to non-ENSO SSTs existed in the AGCMs, these were seemingly erroneous and served to obscure the true skill source related to ENSO.

A new SST skill source, of apparent non-ENSO origins, was discovered for U.S. surface air temperature. Although the AGCM’s winter-spring skill source was again accounted for by a univariate linear influence of ENSO, the AGCM skill significantly exceeded that of either the linear models during the period from late summer through Fall. During these latter seasons, both the AGCMs’ U.S. averaged correlation skill and the spatial extent of its significant local correlations greatly exceeded those occurring in the reduced phase space linear model simulations based on tropical SSTs alone. Our investigation focused on the November-January season of maximum difference in skills, and a coherent source related to SST variability over the subtropical west Pacific Ocean and the South China Sea was found.

That this empirical relation constituted a cause-effect link was confirmed through analysis of an additional suite of AGCM experiments that used an idealization of SST forcing confined to the far west Pacific/South China Sea region. Those runs yielded a pattern of U.S., surface warming in response to warm SST forcing that reproduced the characteristic correlation patterns derived

from both the globally forced multi-model ensemble and the observational data sets. A linear bivariate regression model using only indices of Nino3.4 and west Pacific/South China Sea SSTs explained virtually all of the AGCM simulation skill during the NDJ season.

Hindcast experiments for 1951-1999 were generated in order to determine what elements of skillfully simulated U.S. responses to known SST forcing could be predicted. These employed the multivariate regression model using the 1-season lag relationships between tropical SST predictors and the AGCM's U.S. seasonal climate. These are analogous to zero-lead predictions, and estimate what the AGCM predictions would have been given a forecast for the tropical SSTs. Virtually all of the simulation skill was recovered in the hindcast experiments for the seasons and the variables for which ENSO was the dominant source of simulation skill. The hindcasts failed to recover the AGCM simulation skill of U.S. surface air temperature during late summer-fall. It was demonstrated, however, that a 1-season lagged bivariate model could recover the AGCM's high simulation skill during fall. One conclusion drawn from the hindcasts is that the principal SST sources of U.S. predictability can themselves be skillfully forecast at zero-lead.

b. Discussion

Dynamical methods for U.S. climate prediction are now achieving parity with empirical approaches (e.g., Peng et al. 2000). The key question is whether such equivalence of capability testifies to the limitations in skill of seasonal climate prediction, as suggested by Anderson et al. (1999), or whether it speaks to the infancy of dynamical models. Our diagnosis of the skill sources offers the following insight on such questions.

The fact that the linear response to ENSO SST variability accounts for much of the AGCM skill over the U.S. ---- particularly in winter/spring for both precipitation and temperature---- implies that a simple regression model trained from observations should have comparable skill to the dynamical model given sufficient data. We confirm this to be the case in Fig. 12 where the bar graphs compare the hindcast skill of our prior multivariate model to an identical model in which the U.S. predictand data are derived from observations. Their 50-year hindcast skills are indistinguishable, and the reason is because both are originating from the ENSO influence whose response is realistically modeled in the AGCMs. If there are other tropical SST skill sources in nature, then they are evidently too small (or they occur too infrequently) to be detected and have material effects on the 50-year *averaged* skill (This does not discount the possible skillful contribution of such additional tropical sources in individual years). The preeminence of ENSO as a U.S. skill source is further verified by the result of our singular value decomposition (SVD) analysis. We found that the singular value of the leading SVD mode is much larger than other SVD modes. Our result is consistent with that predicted by DelSole and Chang (2003). Our study did identify an additional non-ENSO skill contributor, but as was the case with ENSO, a linear empirical model could be used to capitalize upon that source too.

In this study, our analysis has been focused on the multi-model ensemble mean. Treating predictions from different models as a single ensemble increases the ensemble size and reduces sampling biases. Furthermore the skill of the ensemble mean is found to be higher than the skill of individual AGCMs included (figure not shown). In operational practice, the skill of seasonal forecasts may be further raised by applying more sophisticated multi-model combination

strategies (e.g. Goddard et al. 2003, Barnston et al. 2004). The sources of skill identified in our analysis are limited only to those that can be captured by the multi-model ensemble mean. Additional sources may be identified when optimal ensemble average methods are applied to maximize the skill. It is also possible that the seasonal forecast skill can be further improved, and additional non-ENSO skill sources identified, when newer generation models are available.

The possibility that extratropical SST variations contribute to skill cannot be discounted based upon our analysis. It is possible that the near equivalence in simulation skill between the AGCMs and the statistical models using only tropical SST predictors (e.g., in all seasons for U.S. precipitation, and in winter/spring for U.S. surface air temperature) is symptomatic of AGCM biases, including the two-tier design of experiments, rather than a lack of predictive impact of non-tropical SSTs. A important thrust for future investigation is assessing how, if at all, this current diagnosis of U.S. seasonal forecast skill related to sea surface temperature influences is modified when using the next generation of atmospheric GCMs. Likewise, a comparative analysis of the AGCM skill against that of coupled ocean-atmosphere models is needed to specifically address whether the predictability estimates, such as presented herein and in previous studies using two-tiered systems, are applicable to fully coupled Earth System models. Regarding the skill attributes of such fully coupled models, it should also be noted that initialization of observed land surface conditions may be sources for U.S. skill, in addition to ocean surface conditions. While the AGCM studied herein do employ coupled soil moisture models of various complexity, none of the runs used observed soil moisture conditions.

A further open question in seasonal climate predictability concerns the expected value for skill itself. Is an analysis of skill drawn from 50 years of data a reliable estimate of the true expected value? To what degree does such a 50-year analysis speak to the system's predictability in general, and what are error bars on those estimates owing to sampling alone? Little is known about the manner in which skill can vary over extended time periods. In so far as ENSO has been affirmed to be the primary source of skill, it is reasonable to suspect that periods of low ENSO variance would lead to lower U.S. climate predictability. At the same time, ENSO itself accounts for only a small fraction of U.S. climate variability (e.g., Hoerling and Kumar 2002), and thus purely atmospheric intrinsic variations could also impact skill fluctuations. As a further diagnosis of the statistical robustness of our skill, we have used 650 years of output from an integration of an unforced coupled ocean-atmosphere model (NCAR's CCSM2; Kiehl and Gent 2004). This model possesses ENSO related variability that is sufficiently realistic for our purposes in so far as the model's ENSO signal over the U.S. explains roughly the same fraction of interannual variance as does the observed ENSO signal (not shown). Another empirical multivariate model was developed for 1-season lag relationships between tropical SSTs and U.S. surface air temperature based on the first 150 years of the coupled model run. Hindcasts were then made for the subsequent 500 years of model data. Figure 13 shows the probability densities of independent 50-year samples of hindcast skill of U.S. surface temperature for the seasons JFM through MAM. The 50-yr mean variations range from a low correlation skill of 0.1 and a high value of 0.3, with a median value near 0.2. These variations arise purely due to intrinsic coupled model noise. This result suggests that dynamical models are still needed to fully harvest the skill source of seasonal forecasts even if the skill of dynamical models comes from just one degree of

freedom in the SST, since the skills of empirical methods trained on 50-year of data may have large decadal swings.

In conclusion, it is apparent that a sound appraisal of the future prospects for U.S. seasonal forecast skill must be founded upon an understanding of the sources for skill itself. Here we have explored the ocean's role, and identified a single degree of freedom in SST variations as providing the bulk of U.S. seasonal skill. An outlook for future capabilities must also seek to understand the sources for multi-decadal variations in skill, which our analysis indicates could occur simply due to low frequency variations in ENSO and other intrinsic fluctuations of the coupled ocean-atmosphere system. If multi-decadal variations in seasonal forecast skill are sufficiently large, then they could mask technological advances in seasonal prediction methodologies.

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References

- Anderson, J. H. van den Dool, A. Barnston, W. Chen, W. Stern, and J. Ploshay, 1999: Present-day capabilities of numerical and statistical models for atmospheric extratropical seasonal simulation and prediction. *Bull. Amer. Meteor. Soc.*, **80**, 1349-1361
- Barnett, T.P. and R.W. Preisendorfer, 1987: Origins and levels of monthly and seasonal forecast skill for North American surface temperature determined by canonical correlation analysis. *Mon. Wea. Rev.*, **115**, 1825-1850
- Barnston, A.G., 1994: Linear statistical short-term climate predictive skill in the northern hemisphere. *J. Climate*, **7**, 1513-1564
- Barnston, A. G. and T.M. Smith, 1996: Specification and prediction of global surface temperature and precipitation from global SST using CCA. *J. Climate*, **9**, 2660-2697
- Barnston, A.G., S.J. Mason, L. Goddard, D.G. Dewitt, and S.E. Zebiak, 2003: Multimodel ensembling in seasonal climate forecasting at IRI. *Bull. Amer. Meteor. Soc.*, **84**, 1783-1796
- Barsugli, J.J. and P.D. Sardeshmukh. 2002: Global atmospheric sensitivity to tropical SST anomalies throughout the Indo-Pacific basin. *J. Climate*, **15**, 3427–3442.
- Broccoli, A.J. and S. Manabe. 1992: The effects of orography on midlatitude northern hemisphere dry climates. *J. Climate*, **5**, 1181–1201.

Chen, P., and M. Newman, 1998: Rossby wave propagation and the rapid development of upper-level anomalous anticyclones during the 1988 U.S. drought. *J. Climate*, **11**, 2491-2504.

Delsole, T. and P. Chang, 2003: Predictable component analysis, canonical correlation analysis, and autoregressive models. *J. Atmos. Sci.*, **60**, 409-416

Geisler, J.E., M.L. Blackmon, G.T. Bates, and S. Muñoz, 1985: Sensitivity of January climate response to the magnitude and position of equatorial Pacific sea surface temperature anomalies. *J. Atmos. Sci.*, **42**, 1037-1049

Goddard, L. and S.J. Mason, 2002: Sensitivity of seasonal climate forecasts to persisted SST anomalies. *Clim. Dyn.*, **19**, 619-631

Goddard, L., A.G. Barnston, and S.J. Mason, 2003: Evaluation of the IRI's "net assessment" seasonal climate forecasts 1997-2001. *Bull. Amer. Meteor. Soc.*, **84**, 1761-1780

Held, I. M., S. W. Lyons, and S. Nigam, 1989: Transients and the extratropical response to El Niño. *J. Atmos. Sci.*, **46**, 163-174.

Higgins, R.W., H.-K. Kim and D. Unger. 2004: Long-lead seasonal temperature and precipitation prediction using tropical Pacific SST consolidation forecasts. *J. Climate*, **17**, 3398–3414.

Hoerling, M. P., and A. Kumar, 2002: Atmospheric response patterns associated with tropical forcing. *J. Climate*, **15**, 2184-2203.

Hoerling, M. P., and A. Kumar, 2003: The perfect ocean for drought. *Science*, **299**, 691-694.

Hoskins, B.J. and D.J. Karoly, 1981: The steady linear response of a spherical atmosphere to thermal and orographic forcing. *J. Atmos. Sci.*, **38**, 1179-1196

Kiehl, J.T., J. J. Hack, G. B. Bonan, B. A. Boville, D. L. Williamson and P. J. Rasch. 1998: The National Center for Atmospheric Research community climate model: CCM3. *J. Climate*, **11**, 1131–1149.

Kiehl, J. T., and P. R. Gent , 2004: The community climate system model, version 2. *J. Climate*, **17**, 3666–3682

Kiladis, G.N. and H.F. Diaz, 1989: Global climatic anomalies associated with extremes in the southern oscillation. *J. Climate*, **2**, 1069-1090

Kumar A., M. Hoerling, M. Ji, A. Leetmaa and P. Sardeshmukh. 1996: Assessing a GCM's Suitability for Making Seasonal Predictions. *J. Climate*, **9**, 115–129.

Kumar, A., W. Wang, M.P. Hoerling, A. Leetma, and M. Ji, 2001: The sustained North American warming of 1997 and 1998. *J. Climate*, **14**, 345-353

Livezey, R.E. and W.Y. Chen, 1983: Statistical field significance and its determination by Monte

Carlo techniques. *Mon. Wea. Rev.*, **111**, 46-59

Livezey, R.E. and K.C. Mo, 1987: Tropical-extratropical teleconnections during the Northern Hemisphere winter: Part II: Relationships between the monthly mean Northern hemisphere circulation patterns and proxies for tropical convection. *Mon. Wea. Rev.*, **115**, 3115-3132.

North, G.R., T.L. Bell, R.F. Cahalan, and F.J. Moeng, 1982: Sampling errors in the estimation of empirical orthogonal functions. *Mon. Wea. Rev.*, **110**, 699-706

Palmer, T.N. and J.A. Owen. 1986: A possible relationship between some “severe” winters in North America and enhanced convective activity over the tropical west Pacific. *Mon. Wea. Rev.*, **114**, 648–651.

Pegion, P., S. Schubert, and M. J. Suarez, 2000: An Assessment of the Predictability of Northern Winter Seasonal Means with the NSIPP1 AGCM. *NASA Tech. Memo-2000-104606*, Vol. 18, 100pp.

Peng, P., A. Kumar, A.G. Barnston, L. Goddard, 2000: Simulation skills of the SST-forced global climate variability of the NCEP-MRF9 and the Scripps-MPI ECHAM3 models. *J. Climate*, **13**, 3657-3679

Rayner, N. A., D. E. Parker, E. B. Horton, C. K. Folland, L. V. Alexander, D. P. Rowell, E. C. Kent, and A. Kaplan, 2003: Global analyses of sea surface temperature, sea ice, and night marine air temperature since the late nineteenth century. *J. Geophys. Res.*, **108**, 4407,

doi:10.1029/2002JD002670

Roeckner, E., and coauthors. 1992. Simulation of the present-day climate with the ECHAM model: Impact of model physics and resolution. *Max-Planck-Institute für Meteorologie, Report* 93, 171 pp. (Available from MPI für Meteorologie, Bundesstr. 55, D-20146 Hamburg, Germany.)

Ropelewski, C. F. and M.S. Halpert, 1989: Precipitation patterns associated with the high index phase of the southern oscillation. *J. Climate*, **2**, 268-284

Ting, M. F. and L. H. Yu, 1998: Steady response to tropical heating in wavy linear and nonlinear baroclinic models. *J. Atmos. Sci.*, **55**, 3565-3582.

Figure Captions:

Figure 1 Schematic of the empirical model used to simulate and forecast U.S. seasonal precipitation (P) and surface air temperature (Ts). The indices i and j denote the seasons used in each simulation and hindcast experiment. In the simulations, j equals to i for each experiment. In the 1-season-lead hindcasts j is one season lag to i (e.g., when i points to OND of 1998, j is JFM of 1999). See text for further details.

Figure 2 Grids included in the verification for precipitation (all "o" and "+" grid-points), surface temperature (only the "+" grid-points).

Figure 3 Spatial average of the 1951-99 temporal correlation for the US precipitation (top) and field significance values (bottom). Points above the shade are statistically significant at 95% confidence level.

Figure 4 Temporal correlation between observed and simulated/hindcast 3-month mean anomalies of precipitation. Red numbers are the first digit of the correlation coefficient value at each grid-point. Having higher than 90% local significance. Black numbers shown at bottom right corner in each panel indicate field significance level (N/N_0) of the corresponding correlation patterns. Ratios exceeding 1 indicates field significant at 95% confidence level. Note: the "lag-0" in figure caption represents "simulation", and "lag-1" is used for "zero-lead hindcast" in text.

Figure 5 Same as Fig. 3 except for the US surface temperature.

Figure 6 Same as Fig. 4 except for the US surface temperature.

Figure 7 Spatial distribution of the correlation coefficient between SST and an index of regional mean surface temperature over the northern U.S. (42N- 52N,120W-70W) based on AGCM simulated (top panel) and observed (bottom panel) NDJ surface temperatures. Contour interval 0.1. Green box denotes index region for constructing land temperature time series, and white box index region for constructing SST time series.

Figure 8 Time series of NDJ anomalies of Northern U.S. surface temperature (black line) based on AGCM simulations (top) and observations (bottom). Red line denotes the time series of NDJ SSTs averaged over the Western Pacific region. Inserted boxes of Fig. 7 show the averaging domains for constructing each time series.

Figure 9 The 1950-1999 temporal correlation between the NDJ SST anomalies and the NDJ index of subtropical west Pacific SSTs. Contour interval 0.1. See Fig. 7 for the averaging domain for constructing the index. The irregular white contour encloses a region of SST variations coherent with the index region, and is used as the forcing region for idealized AGCM simulations.

Figure 10 Spatial distribution of the correlation between the NDJ SST anomalies in the western Pacific Ocean and the multi-AGCM ensemble mean (top panel) and observed

(middle panel) surface temperature anomalies. Contour interval is 0.1. The CCM3 response to the idealized SST anomalies in the Western Pacific Ocean is shown in the bottom panel. Contour interval is 1. C.

Figure 11 Same as Fig.6 but for the NDJ simulation skill based on a univariate regression model using only an index of subtropical west Pacific (WP) SSTs as predictor, a bivariate regression model using the WP and Nino 3.4 indices as predictors (middle panel); and the hindcast skill based on a lagged bivariate model relating ASO Nino3.4 and WP SST anomalies to NDJ surface temperature (bottom panel),

Figure 12. The U.S. spatially averaged multivariate model hindcast skill for 1951-99 for observed U.S. precipitation, and surface temperature using identical empirical models except that the predicand data are derived from the AGCMs (MR(G)) in one case and from the observations in the other case (MR(O)).

Figure 13. Probability distribution of the 50-year U.S. surface temperature skill obtained from the 500-year simulation of the NCAR coupled climate system model

Table 1. Characteristics of the atmospheric GCMs used in the study.

	SPECTRAL RESOLUTION	SIGMA LAYERS	ENSEMBLE SIZE	REFERENCE
CCM3	T42	18	12	Kiehl et al. (1998)
ECHAM3	T42	18	10	Roeckner et al.(1992)
GFDL	R30	18	12	Broccoli and Manabe (1992)
MRF9	T40	18	12	Kumar et al. (1996)

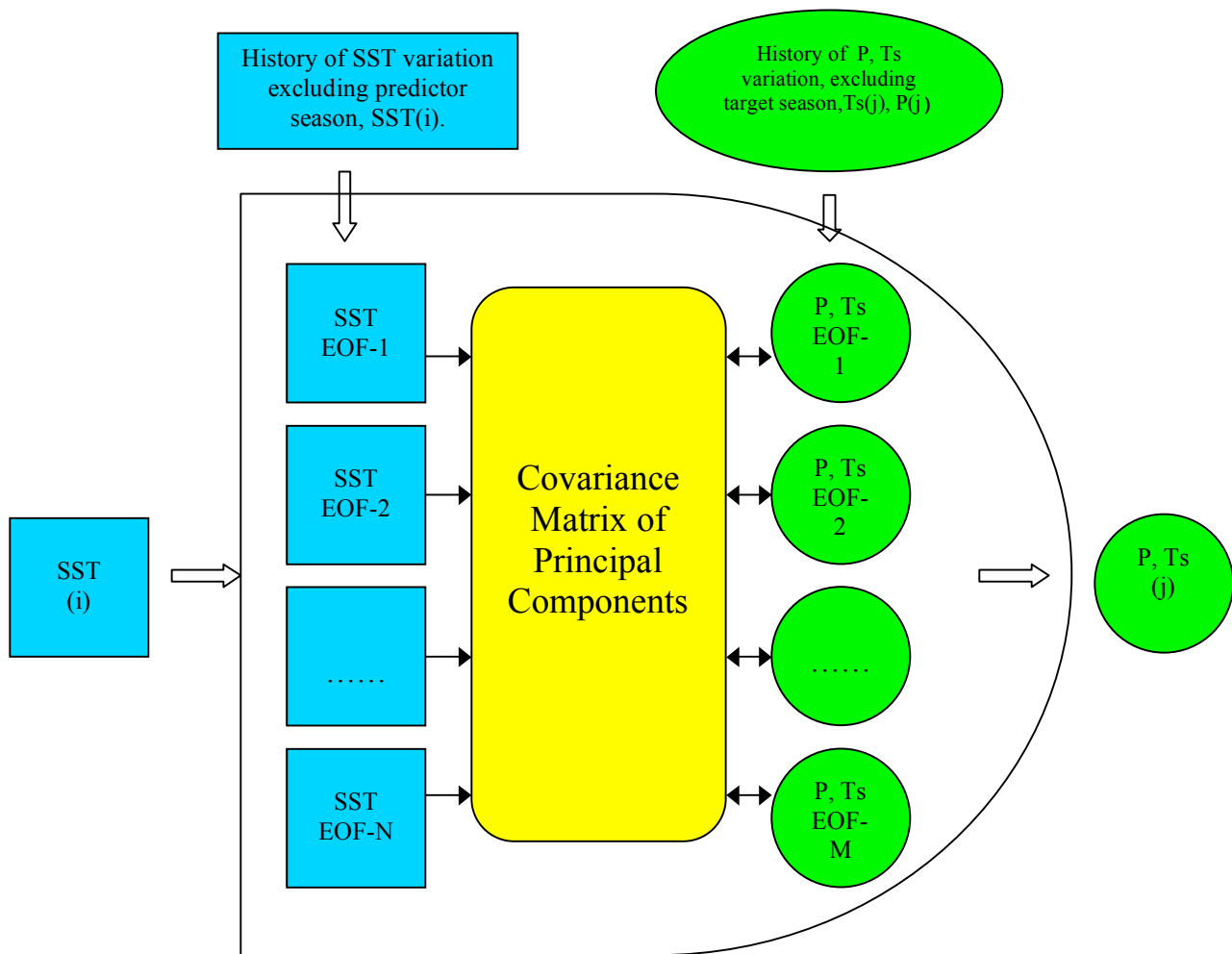


Figure 1 Schematic of the empirical model used to simulate and forecast U.S. seasonal precipitation (P) and surface air temperature (Ts). The indices i and j denote the seasons used in each simulation and hindcast experiment. In the simulations, j equals to i for each experiment. In the 1-season-lead hindcasts j is one season lag to i (e.g., when i points to OND of 1998, j is JFM of 1999). See text for further details.

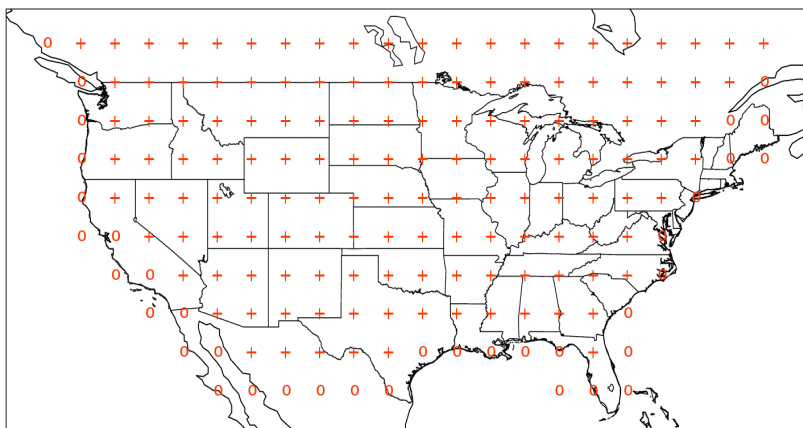


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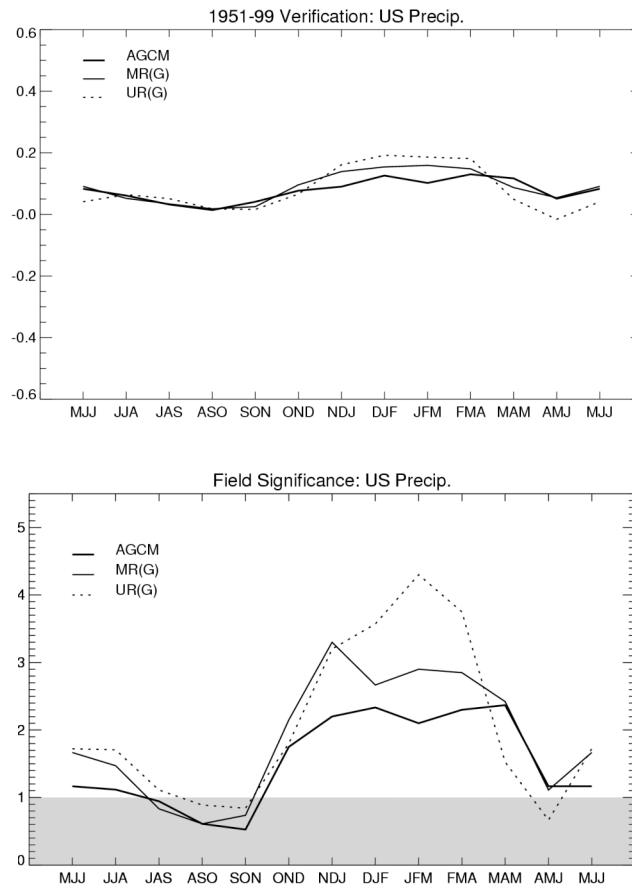


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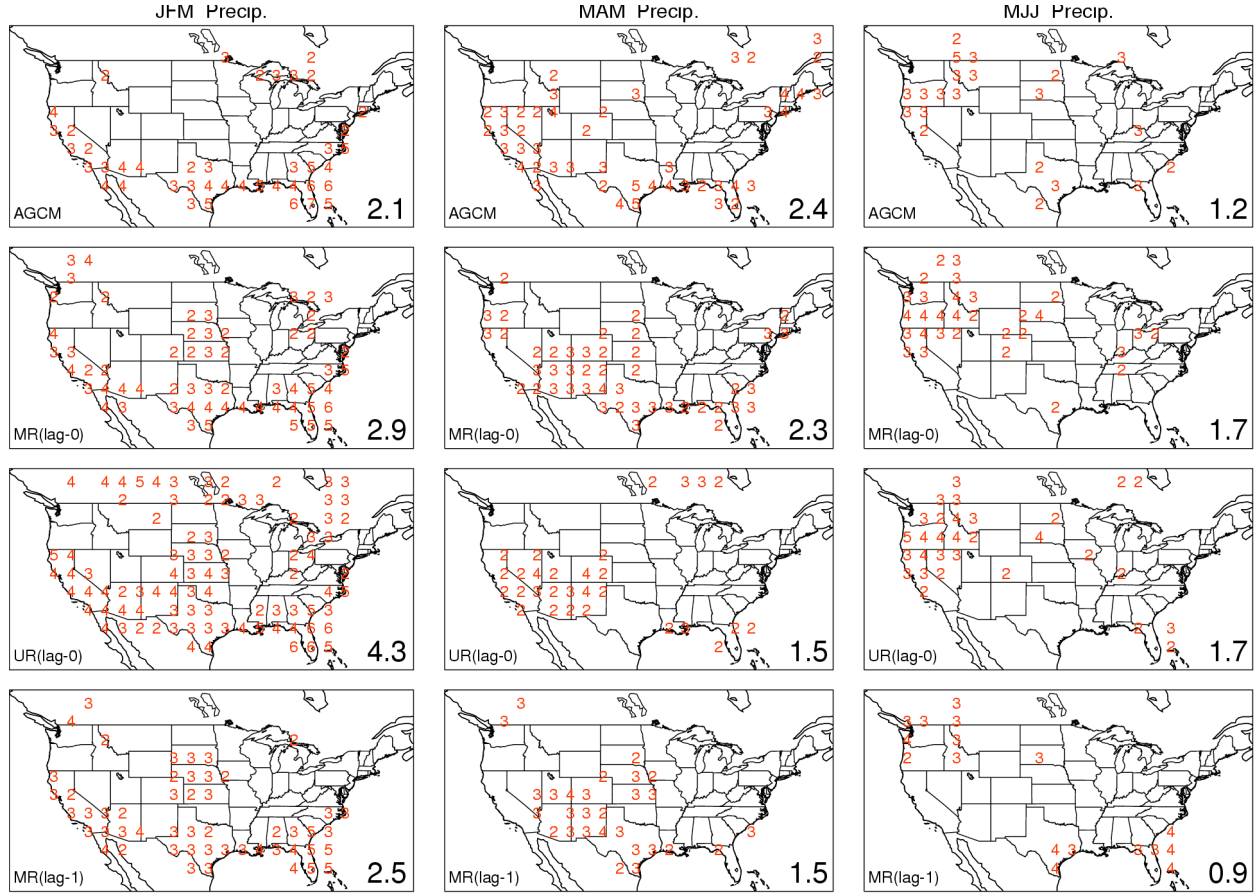


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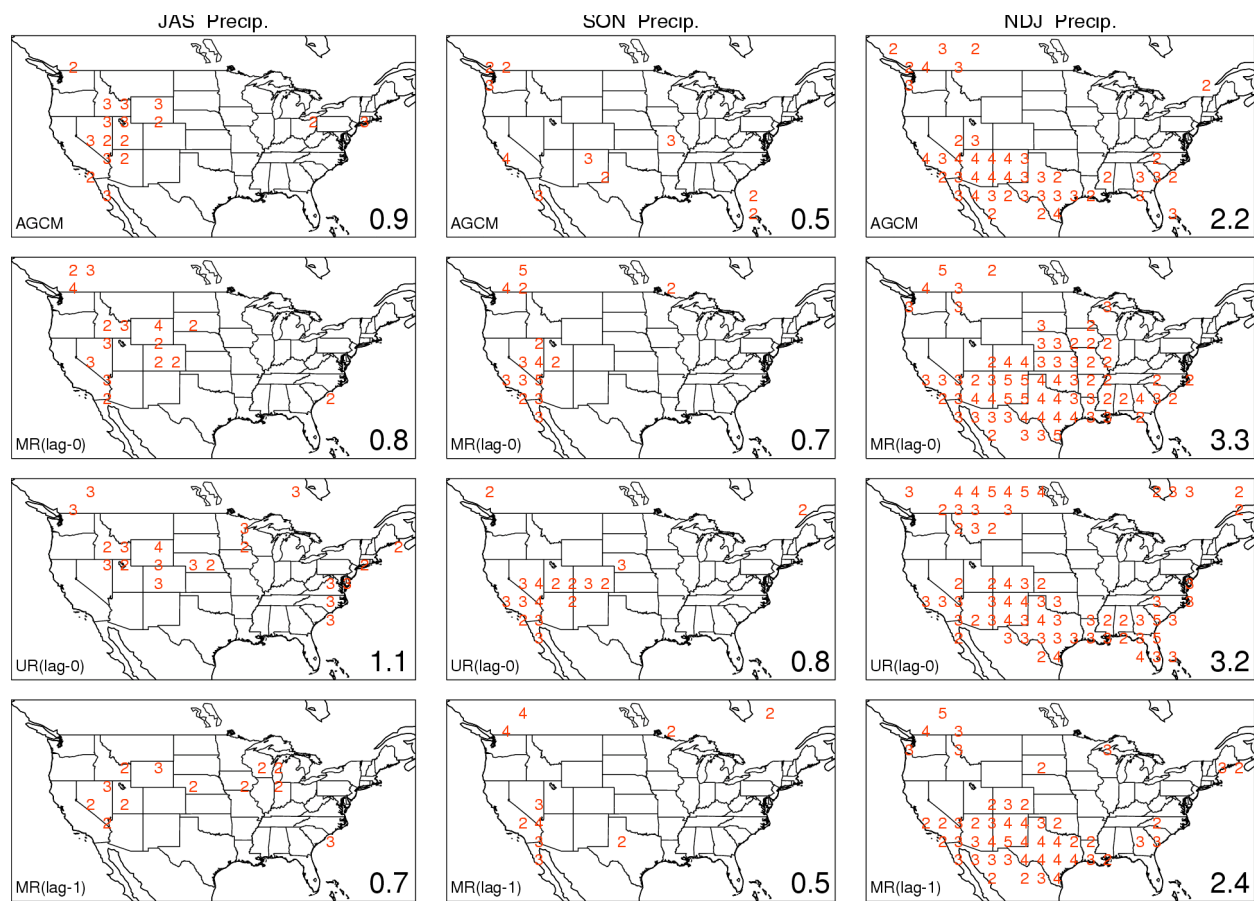


Figure 4 continued.

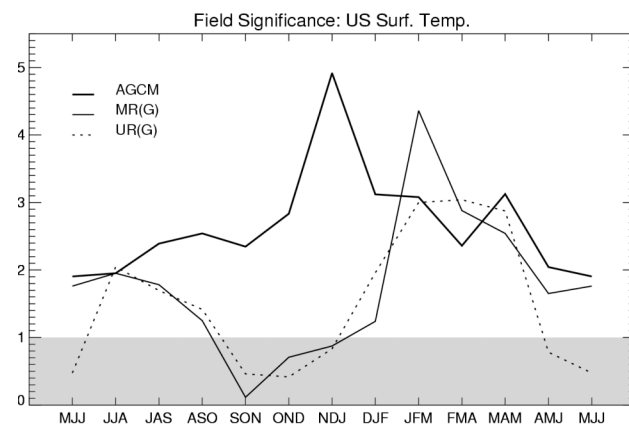
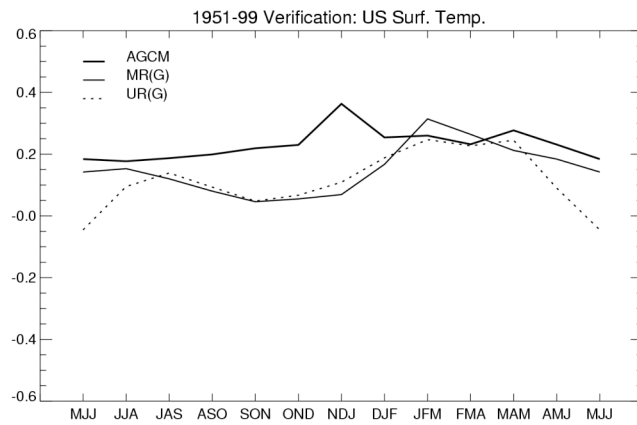


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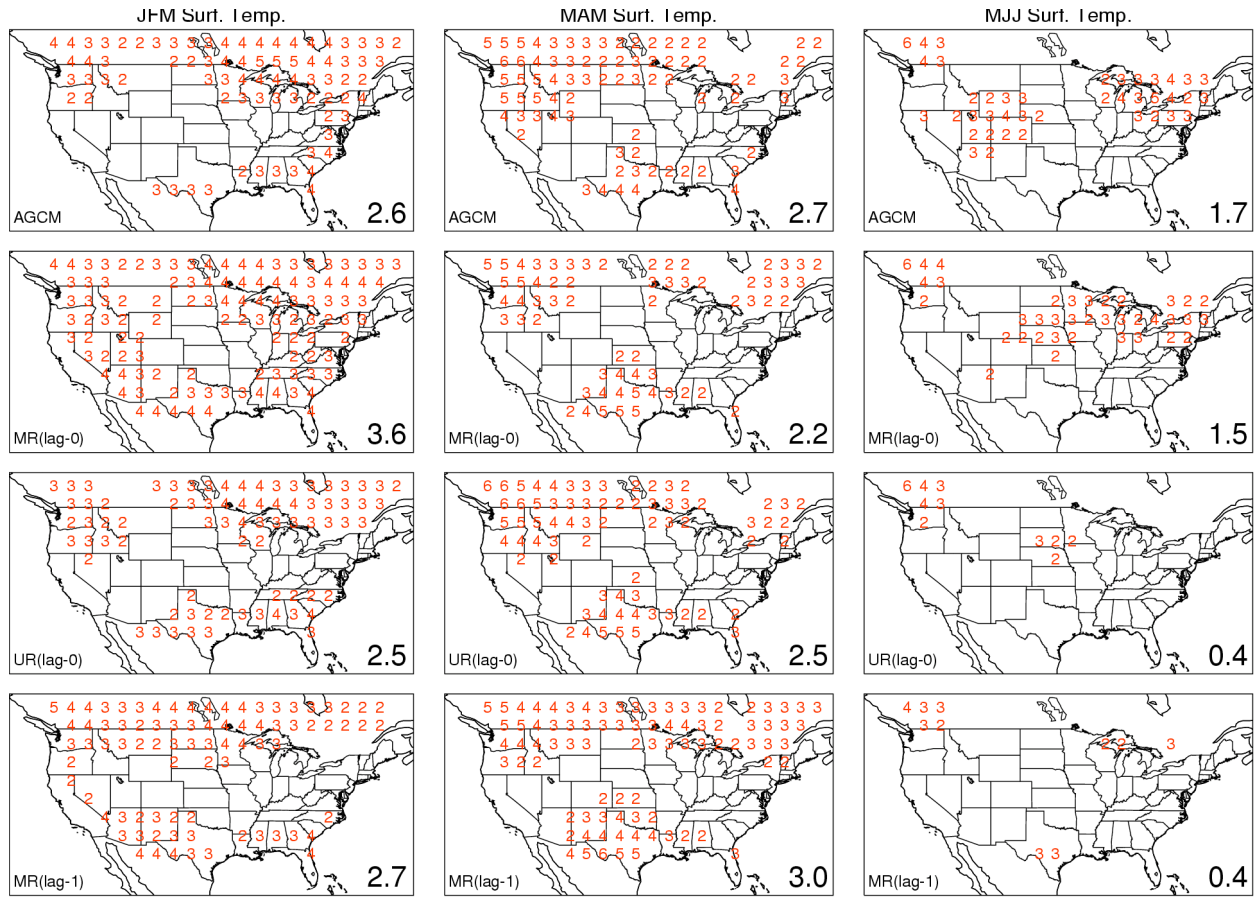


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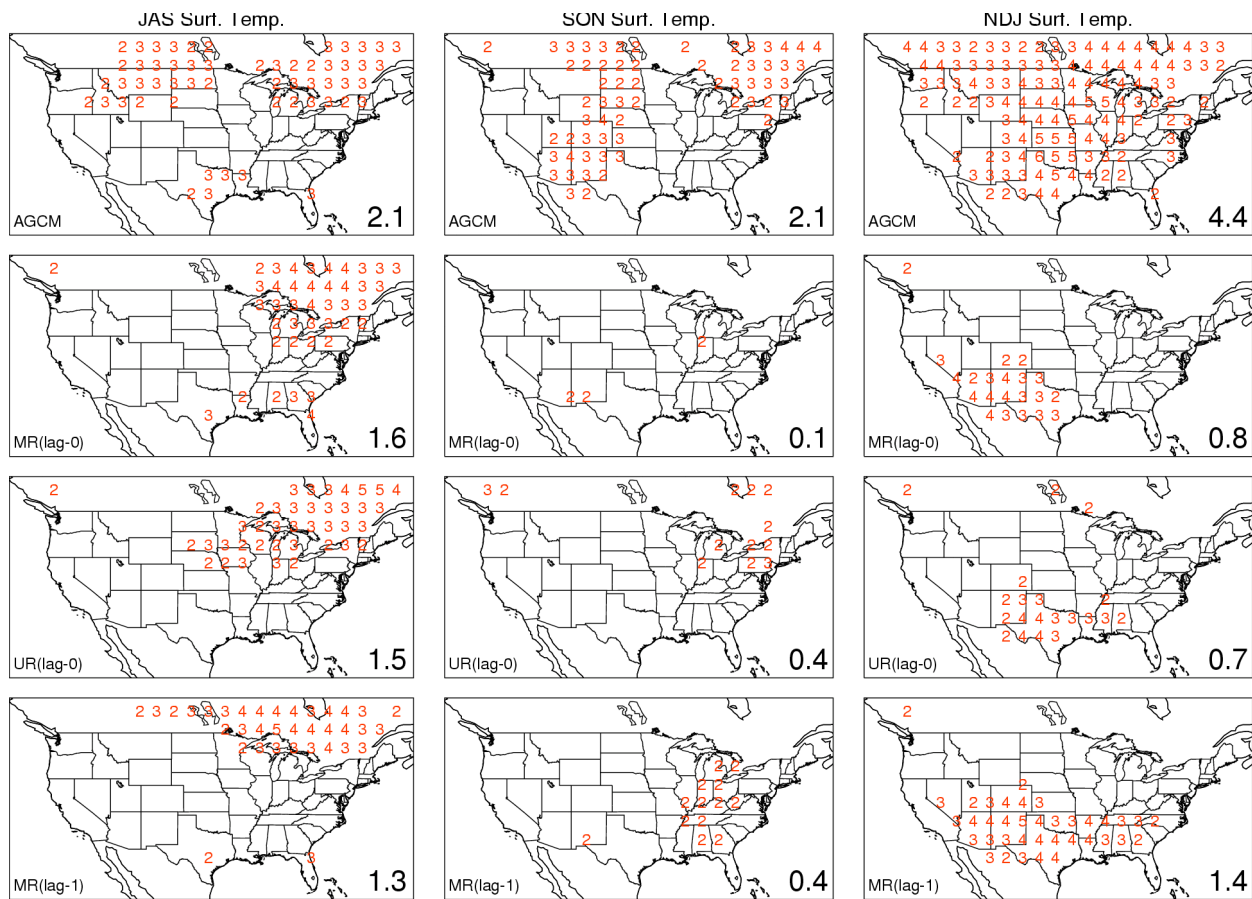


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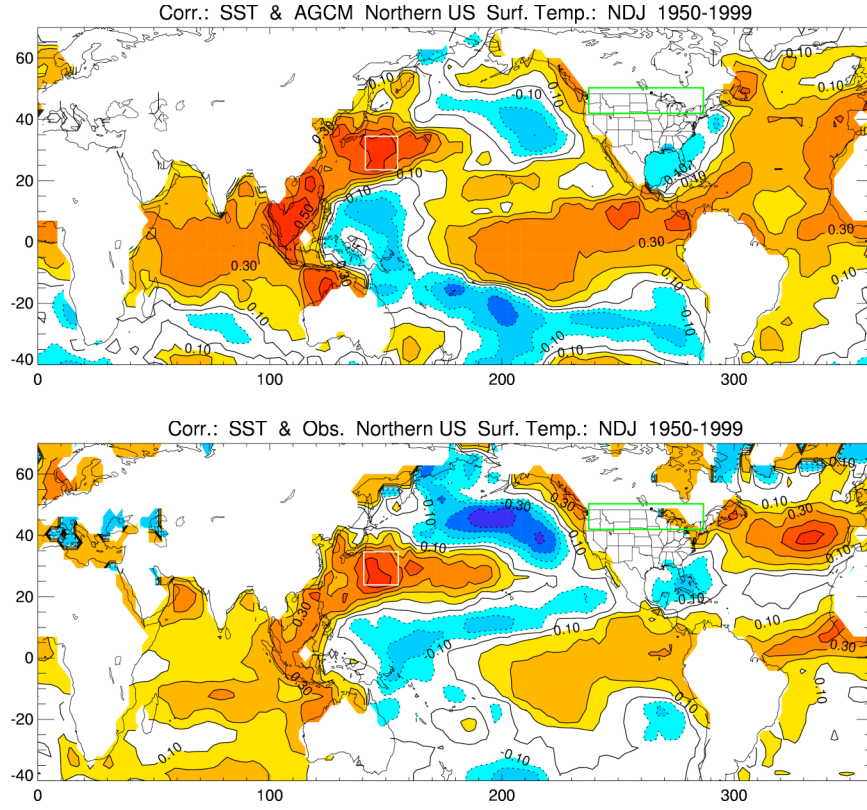


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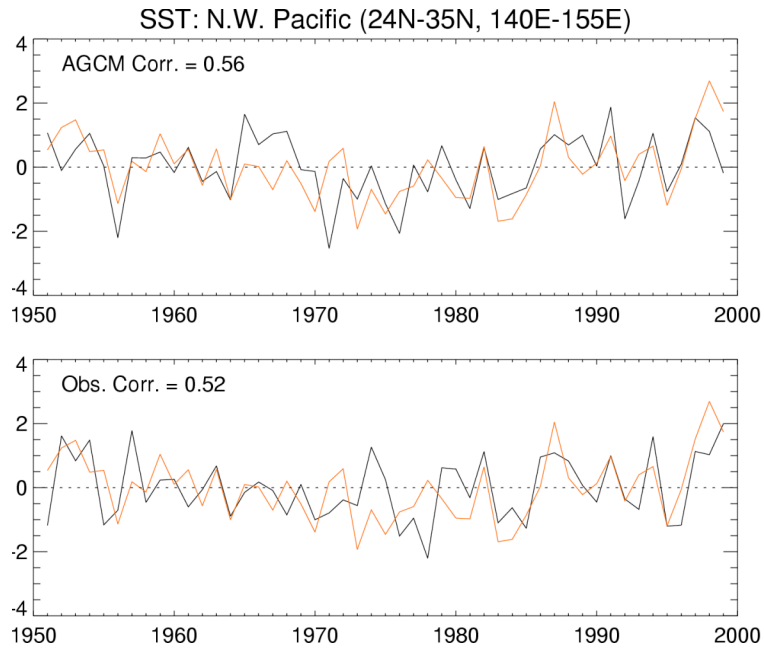


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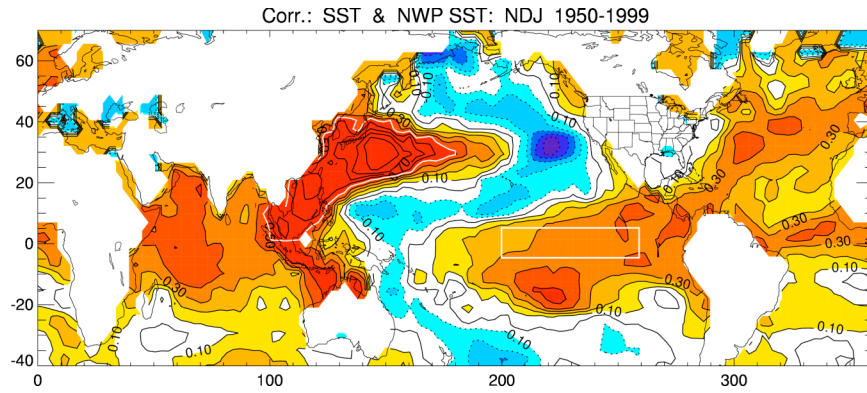


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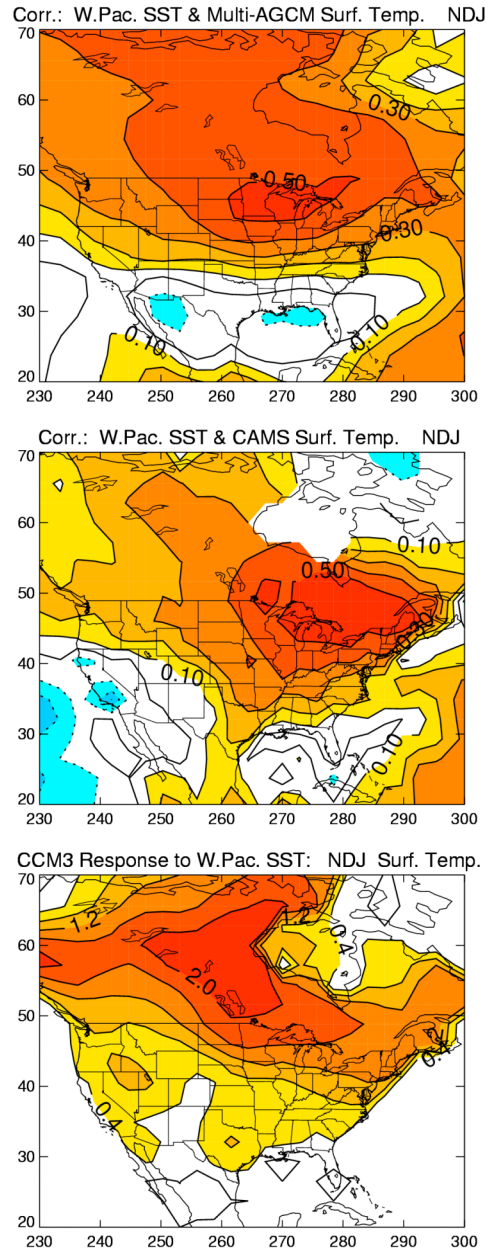


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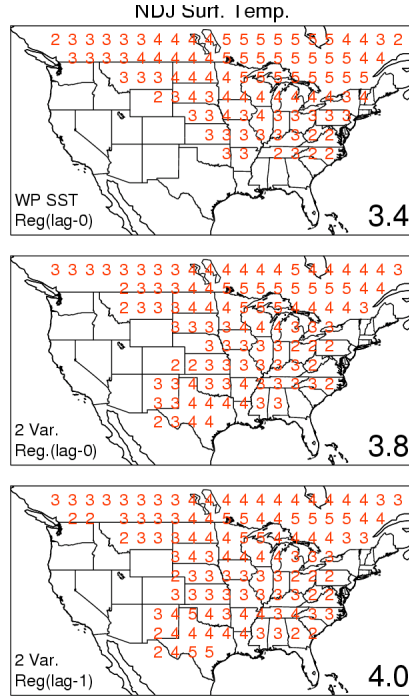


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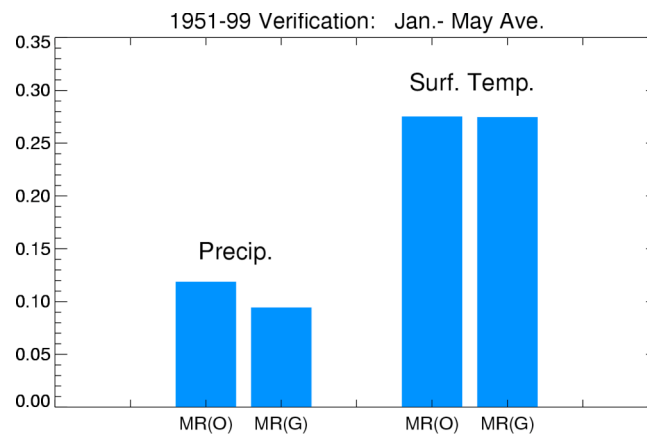


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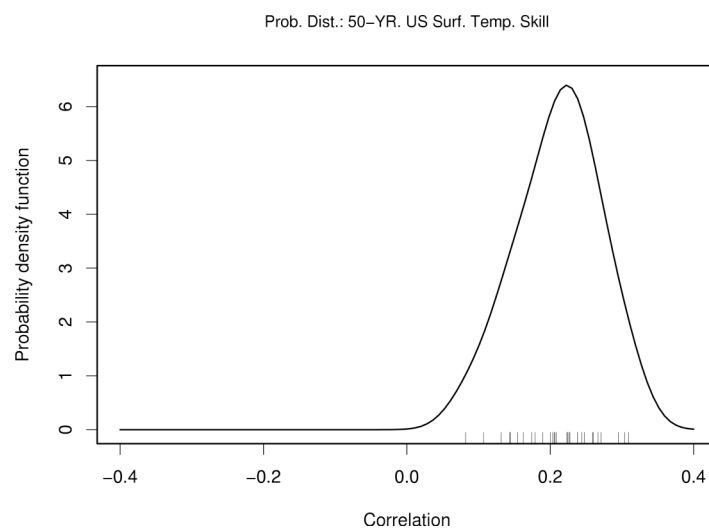


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